Lecture 7: Potential vorticity, Rossby waves and internal waves



Callum J. Shakespeare

Fellow, Climate and Fluid Physics, ANU





2-layer SWE from Lecture 6 $\frac{Du_0}{Dt} - fv_0 = g \frac{\partial \eta_0}{\partial x} \qquad \frac{Du_1}{Dt} - fv_1 = g \frac{\partial \eta_0}{\partial x} + g' \frac{\partial \eta_1}{\partial x}$ $\frac{Dv_0}{Dt} + fu_0 = g \frac{\partial \eta_0}{\partial y} \qquad \frac{Dv_1}{Dt} + fu_1 = g \frac{\partial \eta_0}{\partial y} + g' \frac{\partial \eta_1}{\partial y}$ $\frac{Dh_0}{Dt} = -h_0 \nabla \cdot \mathbf{u_0} \qquad \frac{Dh_1}{Dt} = -h_1 \nabla \cdot \mathbf{u_1}$ $h_0 = \eta_1 - \eta_0 \qquad h_1 = H - \eta_1$



$$-\frac{\partial}{\partial y} \frac{Du_1}{Dt} - fv_1 = g \frac{\partial \eta_0}{\partial x} + g' \frac{\partial \eta_1}{\partial x}$$
$$\frac{\partial}{\partial x} \frac{Dv_1}{Dt} + fu_1 = g \frac{\partial \eta_0}{\partial y} + g' \frac{\partial \eta_1}{\partial y}$$
$$\frac{Dh_1}{Dt} = -h_1 \nabla \cdot \mathbf{u_1}$$
$$h_1 = H - \eta_1$$



2-layer SWE from Lecture 6 atmosphere $-\frac{\partial}{\partial v}\frac{Du_1}{Dt} - fv_1 = g\frac{\partial\eta_0}{\partial x} + g\frac{\partial\eta_1}{\partial x}$ -Z = ~N, 3. -240 $\frac{\partial}{\partial x} \frac{Dv_1}{Dt} + fu_1 = g \frac{\partial \eta_0}{\partial y} + g \frac{\partial \eta_1}{\partial y}$ 2=-M $\frac{Dh_1}{Dt} = -h_1 \nabla \cdot \mathbf{u_1}$ $\rightarrow \underline{n}$ $h_1 = H - \eta_1$ $\frac{\partial}{\partial x}\left(\frac{\partial v_{1}}{\partial t}+\underline{u}_{1}^{\circ}\nabla v_{1}+fv_{1}\right)-\frac{\partial}{\partial y}\left(\frac{\partial u_{1}}{\partial t}+\underline{u}_{1}^{\circ}\nabla u_{1}-fv_{1}\right)=0.$ $\Rightarrow \left(\underbrace{\exists t}_{+} \underbrace{\forall}_{0} \underbrace{\nabla}\right) \left(\underbrace{\exists v_{1}}_{\exists v_{1}} - \underbrace{\exists v_{1}}_{\forall v_{1}}\right) + \underbrace{\exists x_{1}}_{\exists v_{1}} \underbrace{\forall v_{1}}_{\forall v_{1}} - \underbrace{\exists y_{1}}_{\forall v_{1}} \underbrace{\forall v_{1}}_{\forall v_{1}} \underbrace{\forall v_{1}}$ $+f\left(\frac{\partial v_1}{\partial z}+\frac{\partial v_1}{\partial y}\right)+v\frac{\partial f}{\partial y}=0.$

2-layer SWE from Lecture 6 atmosphere $-\frac{\partial}{\partial v}\frac{Du_1}{Dt} - fv_1 = g\frac{\partial\eta_0}{\partial x} + g\frac{\partial\eta_1}{\partial x}$ -Z = ~N, 3. -240 $\frac{\partial}{\partial x} \frac{Dv_1}{Dt} + fu_1 = g \frac{\partial \eta_0}{\partial y} + g \frac{\partial \eta_1}{\partial y}$ 2=-M, $\frac{Dh_1}{Dt} = -h_1 \nabla \cdot \mathbf{u_1}$ $\rightarrow \underline{N}$ $h_1 = H - \eta_1$ $\frac{\partial}{\partial x}\left(\frac{\partial v_{1}}{\partial t}+\underline{u}_{1}^{\circ}\nabla v_{1}+fv_{1}\right)-\frac{\partial}{\partial y}\left(\frac{\partial u_{1}}{\partial t}+\underline{u}_{1}^{\circ}\nabla v_{1}-fv_{1}\right)=0.$ $\Rightarrow \left(\underbrace{\partial}_{\partial t} + \underbrace{\mathcal{Y}}_{1}, \circ \nabla \right) \left(\underbrace{\partial \mathcal{V}}_{\partial z} - \underbrace{\partial \mathcal{V}}_{\partial y} \right) + \underbrace{\partial \mathcal{Y}}_{\partial z} \cdot \nabla \mathcal{V}_{1} - \underbrace{\partial \underline{\mathcal{Y}}}_{\partial y} \cdot \nabla \mathcal{V}_{1} \right)$ $\underbrace{\partial}_{Dt} + f\left(\underbrace{\partial \mathcal{V}}_{\partial z} + \underbrace{\partial \mathcal{V}}_{\partial y} \right) + \sqrt{\frac{\partial f}{\partial y}} = 0.$



$$-\frac{\partial}{\partial y} \frac{Du_1}{Dt} - fv_1 = g \frac{\partial \eta_0}{\partial x} + g \frac{\partial \eta_1}{\partial x}$$
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$$\frac{Dh_1}{Dt} = -h_1 \nabla \cdot \mathbf{u}_1$$
$$h_1 = H - \eta_1$$



$$\frac{D\omega_{i}}{Dt} + \omega_{i}\nabla \cdot \underline{u}_{i} + f\nabla \cdot \underline{u}_{i} + v\frac{\delta f}{\delta y} = 0$$

$$\implies \frac{D}{Dt}(f + \omega_{i}) + (\omega_{i} + f)(-\frac{1}{h_{i}}\frac{Dh_{i}}{Dt}) = 0$$

$$\implies \frac{D}{Dt}\left(\frac{\omega_{i} + f}{h_{i}}\right) = 0$$

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$$\frac{D\omega_{1}}{Dt} + \omega_{1}\nabla_{2}\underline{u}_{1} + f\nabla_{2}\underline{u}_{1} + v\frac{\delta f}{\delta y} = 0$$

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$$\implies \frac{D}{Dt}\left(\frac{\omega_{1}+f}{h_{1}}\right) = 0$$

The quantity

$$q_1 = \frac{\omega_1 + f}{h_1} = \frac{\omega_1 + f}{H - \eta_1}$$
is conserved
This is the **Potential Vorticity**

$$-\frac{\partial}{\partial y} \frac{Du_1}{Dt} - fv_1 = g \frac{\partial \eta_0}{\partial x} + g \frac{\partial \eta_1}{\partial x}$$
$$\frac{\partial}{\partial x} \frac{Dv_1}{Dt} + fu_1 = g \frac{\partial \eta_0}{\partial y} + g \frac{\partial \eta_1}{\partial y}$$
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This is the **Potential Vorticity**
Same reasoning for layer 0:

$$q_{0} = \frac{\omega_{0} + f}{h_{0}} = \frac{\omega_{0} + f}{\eta_{1} - \eta_{0}}$$
is conserved

$$-\frac{\partial}{\partial y} \frac{Du_1}{Dt} - fv_1 = g \frac{\partial \eta_0}{\partial x} + g \frac{\partial \eta_1}{\partial x}$$
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$$q_{0} = \frac{\omega_{0} + f}{h_{0}} = \frac{\omega_{0} + f}{\eta_{1} - \eta_{0}}$$
is conserved

Suppose that: $\omega \ll f~~{\rm and}~~\eta_i \ll H~~{\rm and}$ the flow is STEADY

Then: $q_1 \simeq \frac{f}{H}$

And:
$$\frac{Dq_1}{Dt} = \vec{u} \cdot \nabla q_1 = \vec{u} \cdot \nabla \frac{f}{H} = 0$$



$$q_{1} = \frac{\omega_{1} + f}{h_{1}} = \frac{\omega_{1} + f}{H - \eta_{1}}$$

is conserved
This is the **Potential Vorticity**
Same reasoning for layer 0:
$$q_{0} = \frac{\omega_{0} + f}{h_{0}} = \frac{\omega_{0} + f}{\eta_{1} - \eta_{0}}$$

is conserved

Suppose that: $\omega \ll f$ and $\eta_i \ll H$ and the flow is STEADY

Then: $q_1 \simeq \frac{f}{H}$

And:
$$\frac{Dq_1}{Dt} = \vec{u} \cdot \nabla q_1 = \vec{u} \cdot \nabla \frac{f}{H} = 0$$

The flow follows contours of f/H



Which hemisphere is this in?





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Then: $q_1 \simeq \frac{f}{H}$

And:
$$\frac{Dq_1}{Dt} = \vec{u} \cdot \nabla q_1 = \vec{u} \cdot \nabla \frac{f}{H} = 0$$

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The flow follows contours of f/H







Suppose that: $\omega \ll f$ and $\eta_i \ll H$ and the flow is STEADY

Then: $q_1 \simeq \frac{f}{H}$

And:
$$\frac{Dq_1}{Dt} = \overrightarrow{u_1} \cdot \nabla q_1 = \overrightarrow{u_1} \cdot \nabla \frac{f}{H} = 0$$

The flow follows contours of f/H If H gets smaller, so must |f|....











New 2 layer SW Equations

X momentum

Y momentum

Conserve PV

$$\begin{aligned} \frac{Du_0}{Dt} - fv_0 &= g \frac{\partial \eta_0}{\partial x} & \frac{Du_1}{Dt} - fv_1 &= g \frac{\partial \eta_0}{\partial x} + g' \frac{\partial \eta_1}{\partial x} \\ \frac{Dv_0}{Dt} + fu_0 &= g \frac{\partial \eta_0}{\partial y} & \frac{Dv_1}{Dt} + fu_1 &= g \frac{\partial \eta_0}{\partial y} + g' \frac{\partial \eta_1}{\partial y} \\ \frac{D}{Dt} \left(\frac{\omega_0 + f}{h_0} \right) &= 0 & \frac{D}{Dt} \left(\frac{\omega_1 + f}{h_1} \right) &= 0 \\ h_0 &= \eta_1 - \eta_0 & h_1 &= H - \eta_1 \end{aligned}$$















$$\begin{pmatrix} \frac{\partial}{\partial t} + \bar{U}\frac{\partial}{\partial x} \end{pmatrix} u_0 - f_0 v_0 = -g \frac{\partial \eta'_0}{\partial x} \\ \left(\frac{\partial}{\partial t} + \bar{U}\frac{\partial}{\partial x} \right) v_0 + f_0 u_0 = -g \frac{\partial \eta'_0}{\partial y} \\ \left(\frac{\partial}{\partial t} + \bar{U}\frac{\partial}{\partial x} \right) \left(\omega_0 - \frac{f_0}{H}\eta'_0\right) + \beta v_0 = 0$$



 $\overline{\bigcup} \stackrel{\geq}{\underset{\forall}{\Rightarrow}} < < \widehat{f} \quad \stackrel{\geq}{\underset{\forall}{\Rightarrow}} < < \widehat{f} \quad \text{No timescales of O(f) = no BGWs} \\ or other gravity waves \\ \left(\frac{\partial}{\partial t} + \overline{U}\frac{\partial}{\partial x}\right) u_0 - f_0 v_0 = -g \frac{\partial \eta'_0}{\partial x} \\ \left(\frac{\partial}{\partial t} + \overline{U}\frac{\partial}{\partial x}\right) v_0 + f_0 u_0 = -g \frac{\partial \eta'_0}{\partial y} \\ \left(\frac{\partial}{\partial t} + \overline{U}\frac{\partial}{\partial x}\right) \left(\omega_0 - \frac{f_0}{H}\eta'_0\right) + \beta v_0 = 0 \quad \begin{array}{c} \text{Keep here as it's} \\ \text{the same order} \\ \text{as } \beta \end{array}$

- This assumption means we are describing "balanced flow" since velocities are in geostrophic balance.
- State known as quasigeostrophic balance (QG)
- There is a more general nonlinear form which we won't look at here...





Linearised barotropic "quasi-geostrophy"

$$\left(\frac{\partial}{\partial t} + \bar{U}\frac{\partial}{\partial x}\right) \left(\frac{g}{f_0}\nabla^2 \eta_0' - \frac{f_0}{H}\eta_0'\right) + \frac{g\beta}{f_0}\frac{\partial \eta_0'}{\partial x} = 0$$

Linearised barotropic "quasi-geostrophy"

$$\frac{1}{3} = n_{0}^{1}$$

$$\frac{1}{3} = n_{0}^{1} = n_{0$$

Linearised barotropic "quasi-geostrophy"

$$\frac{1}{3} = \frac{1}{3} = \frac{1}$$



Rossby waves

$$c_p = \frac{\omega}{k} = \bar{U} - \frac{\beta}{K^2 + \frac{f_0^2}{gH}}.$$

ω/k (km/year)



Wavelength? Distance per year?

Do they match?





Stationary Rossby waves



What about if a flow exists such that $c_p = 0$????



Stationary Rossby waves



What about if a flow exists such that $c_p = 0$????



- A westward travelling wave of 200-300km wavelength would be stationary in an eastward flow of a few cm/s
- What is U is faster than this?

Stationary/trapped Rossby waves



Belonenko et al., 2020









Decouples layer 1 (interior) from layer 0 (surface)

Internal waves



Linearised layer 1 equations with a rigid lid

$$\left(\frac{\partial}{\partial t} + \bar{U}\frac{\partial}{\partial x}\right)u_1 - f_0v_1 = -g'\frac{\partial\eta'_1}{\partial x} \left(\frac{\partial}{\partial t} + \bar{U}\frac{\partial}{\partial x}\right)v_1 + f_0u_1 = -g'\frac{\partial\eta'_1}{\partial y} \left(\frac{\partial}{\partial t} + \bar{U}\frac{\partial}{\partial x}\right)\left(\omega_1 - \frac{f_0}{h_1}\eta'_1\right) + \beta v_1 = 0$$

Internal waves





Linearised layer 1 equations with a rigid lid

$$\left(\frac{\partial}{\partial t} + \bar{U}\frac{\partial}{\partial x}\right)u_1 - f_0v_1 = -g'\frac{\partial\eta'_1}{\partial x} \left(\frac{\partial}{\partial t} + \bar{U}\frac{\partial}{\partial x}\right)v_1 + f_0u_1 = -g'\frac{\partial\eta'_1}{\partial y} \left(\frac{\partial}{\partial t} + \bar{U}\frac{\partial}{\partial x}\right)\left(\omega_1 - \frac{f_0}{h_1}\eta'_1\right) + \beta v_1 = 0$$

Compare with the barotropic equations from before

$$\left(\frac{\partial}{\partial t} + \bar{U}\frac{\partial}{\partial x}\right)u_0 - f_0v_0 = -g\frac{\partial\eta'_0}{\partial x} \left(\frac{\partial}{\partial t} + \bar{U}\frac{\partial}{\partial x}\right)v_0 + f_0u_0 = -g\frac{\partial\eta'_0}{\partial y} \left(\frac{\partial}{\partial t} + \bar{U}\frac{\partial}{\partial x}\right)\left(\omega_0 - \frac{f_0}{H}\eta'_0\right) + \beta v_0 = 0$$

They're the same equations! The internal waves are the same as the surface waves, but with $g \rightarrow g'$ and $H \rightarrow h_1$

Internal Rossby waves





Linearised layer 1 equations with a rigid lid

$$\begin{aligned} \left(\frac{\partial}{\partial t} + \bar{U}\frac{\partial}{\partial x}\right)u_1 - f_0v_1 &= -g'\frac{\partial\eta'_1}{\partial x} \\ \left(\frac{\partial}{\partial t} + \bar{U}\frac{\partial}{\partial x}\right)v_1 + f_0u_1 &= -g'\frac{\partial\eta'_1}{\partial y} \\ \left(\frac{\partial}{\partial t} + \bar{U}\frac{\partial}{\partial x}\right)\left(\omega_1 - \frac{f_0}{h_1}\eta'_1\right) + \beta v_1 &= 0 \end{aligned}$$
$$\begin{aligned} \frac{\partial}{\partial t} + \bar{U}\frac{\partial}{\partial x}\right)\left(\frac{g'}{f_0}\nabla^2\eta'_1 - \frac{f_0}{h_1}\eta'_1\right) + \frac{g'\beta}{f_0}\frac{\partial\eta'_1}{\partial x} = 0 \end{aligned}$$
$$c_p &= \frac{\omega}{k} = \bar{U} - \frac{\beta}{K^2 + \frac{f_0^2}{g'h_1}}. \end{aligned}$$

Internal Rossby waves are SLOWER For a given flow speed, trapped waves are longer



Linearised layer 1 equations with a rigid lid

$$\left(\frac{\partial}{\partial t} + \bar{U}\frac{\partial}{\partial x}\right)u_1 - f_0v_1 = -g'\frac{\partial\eta'_1}{\partial x} \left(\frac{\partial}{\partial t} + \bar{U}\frac{\partial}{\partial x}\right)v_1 + f_0u_1 = -g'\frac{\partial\eta'_1}{\partial y} \left(\frac{\partial}{\partial t} + \bar{U}\frac{\partial}{\partial x}\right)\left(\omega_1 - \frac{f_0}{h_1}\eta'_1\right) + \beta v_1 = 0$$



Linearised layer 1 equations with a rigid lid

$$\left(\frac{\partial}{\partial t} + \bar{U}\frac{\partial}{\partial x}\right)u_{1} - f_{0}v_{1} = -g'\frac{\partial\eta'_{1}}{\partial x}$$

$$\left(\frac{\partial}{\partial t} + \bar{U}\frac{\partial}{\partial x}\right)v_{1} + f_{0}u_{1} = -g'\frac{\partial\eta'_{1}}{\partial y}$$

$$\left(\frac{\partial}{\partial t} + \bar{U}\frac{\partial}{\partial x}\right)\left(\omega_{1} - \frac{f_{0}}{h_{1}}\eta'_{1}\right) + \beta v_{1} = 0$$

$$\left(\frac{\partial}{\partial t} + \bar{U}\frac{\partial}{\partial x}\right)u_{1} - f_{0}v_{1} = -g'\frac{\partial\eta'_{1}}{\partial x}$$

$$\left(\frac{\partial}{\partial t} + \bar{U}\frac{\partial}{\partial x}\right)u_{1} - f_{0}v_{1} = -g'\frac{\partial\eta'_{1}}{\partial x}$$

$$\left(\frac{\partial}{\partial t} + \bar{U}\frac{\partial}{\partial x}\right)v_1 + f_0u_1 = -g'\frac{\partial\eta_1}{\partial y}$$
$$q_1 = \omega_1 - \frac{f_0}{h_1}\eta_1' = \text{const.}$$

Internal waves: balanced wave PV, varying wave momentum **Rossby waves:** balanced wave momentum, varying wave PV



$$\left(\frac{\partial}{\partial t} + \bar{U}\frac{\partial}{\partial x}\right)u_1 - f_0v_1 = -g'\frac{\partial\eta'_1}{\partial x}$$
$$\left(\frac{\partial}{\partial t} + \bar{U}\frac{\partial}{\partial x}\right)v_1 + f_0u_1 = -g'\frac{\partial\eta'_1}{\partial y}$$
$$q_1 = \omega_1 - \frac{f_0}{h_1}\eta'_1 = \text{const.}$$

$$\begin{array}{l} \text{let} \quad \mathbf{\hat{q}} = \hat{\mathbf{\hat{q}}} e^{i(kn+ly-wt)} \\ (-iw+iku)\hat{\mathbf{u}}_{i} - \hat{\mathbf{f}}_{0}\hat{\mathbf{v}}_{i} = -ikg\hat{\mathbf{n}}_{i} \\ (-iw+iku)\hat{\mathbf{v}}_{i} + \hat{\mathbf{f}}_{0}\hat{\mathbf{u}}_{i} = -ilg\hat{\mathbf{n}}_{i} \\ (ik\hat{\mathbf{v}}_{i}, -il\hat{\mathbf{u}}_{i}, -\frac{f_{0}}{h_{i}}\hat{\mathbf{n}}_{i}' = 0 \end{array}$$



$$\begin{pmatrix} \frac{\partial}{\partial t} + \bar{U}\frac{\partial}{\partial x} \end{pmatrix} u_1 - f_0 v_1 = -g'\frac{\partial \eta'_1}{\partial x} \begin{pmatrix} \frac{\partial}{\partial t} + \bar{U}\frac{\partial}{\partial x} \end{pmatrix} v_1 + f_0 u_1 = -g'\frac{\partial \eta'_1}{\partial y} q_1 = \omega_1 - \frac{f_0}{h_1}\eta'_1 = \text{const.}$$



$$\left(\frac{\partial}{\partial t} + \bar{U}\frac{\partial}{\partial x}\right)u_1 - f_0v_1 = -g'\frac{\partial\eta'_1}{\partial x} \left(\frac{\partial}{\partial t} + \bar{U}\frac{\partial}{\partial x}\right)v_1 + f_0u_1 = -g'\frac{\partial\eta'_1}{\partial y} q_1 = \omega_1 - \frac{f_0}{h_1}\eta'_1 = \text{const.}$$

Let $\phi = \hat{\phi} e^{i(kn+ly-wt)}$ $(-iw+ik\pi)\hat{\mu}_{i} - f_{o}\hat{\nu}_{i} = -ikg\hat{n}_{i}$ $(-iw+ik\pi)\hat{\nu}_{i} + f_{o}\hat{\mu}_{i} = -ilg\hat{n}_{i}$ $ik\hat{v}_{1} - i\hat{l}\hat{u}_{1} - \frac{f_{0}}{h}\hat{v}_{1}' = 0$ Etc, etc, etc, $\omega = k\bar{U} \pm \sqrt{f_0^2 + K^2 g' h_1}$ Internal gravity wave dispersion relation







- The bathymetry induces a z (or y) velocity to the flow, which reduces its buoyancy (or vorticity)
- If the perturbation is slow/weak, the flow remains in/near balance and returns to its original course....

The mechanism behind the waves

Gravity waves



Rossby waves



- The bathymetry induces a z (or y) velocity to the flow, which reduces its buoyancy (or vorticity)
- If the perturbation is slow/weak, the flow remains in/near balance and returns to its original course....

The mechanism behind the waves





Rossby waves



- The bathymetry induces a z (or y) velocity to the flow, which reduces its buoyancy (or vorticity)
- If the perturbation is slow/weak, the flow remains in/near balance and returns to its original course....
- But if the perturbation is **fast/strong** ($\overline{U} \sim \frac{\omega}{k}$), it kicks of an oscillation in the lee of the obstacle...

References

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